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Biogeochemistry of the Amazonian floodplains.

Insights from six end-member mixing models.

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Abstract: The influence of Amazonian floodplains on the hydrological, sedimentary and biogeochemical river budget was investigated along the Vargem Grande – Obidos reach, by applying six mixing models based on variable regional and/or variable hydrological sources. By comparing the output of many different models designed for different purposes, the nature and the magnitude of processes linking water and biogeochemical budgets of the Amazonian floodplains were clarified. This study reveals that most of the chemical baseline of the Amazon River basin is acquired before the studied 2000 km Amazonian reach. However, the tight connection between the hydrograph stage of the River and the chemical signals provides insightful information on the dynamics of its floodplains. The chemical expression of biotic and abiotic processes occurring in the Amazonian floodplains can be more particularly perceived during falling waters. It appears delayed in time compared to the maximum extension of submerged area, because the alternating water circulation polarity (filling vs. emptying) between the main channel and the adjacent floodplains determines delayed emptying of floodplains during falling waters. It results also in a longer time of residence in the hydrograph network which strengthens the rate of transformation of

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transiting materials and solutes. Biotic and biologically-mediated processes tend to accentuate changes in river water chemistry initiated upstream, in each sub-basin, along river corridors, indicating that processes operating downstream prolong those from upstream (e.g. floodplains of the large tributaries). Conversely, the flood wave propagation tends to lessen the seasonal variability as a result of the water storage in the floodplains which admixes waters of distinct origins (in time and space). The morphology of floodplains, determining the deposition and the diagenesis of the sediments as well as the variable extension of submerged areas or the chronology of floodplains storage/emptying, appears to be the main factor controlling the floodplains biogeodynamics. By coupling classical end-member mixing models (providing insight on hydrological source) with a variable regional contribution scheme, relevant information on the biogeochemical budget of the Amazonian floodplains can be achieved.

Keywords: Amazon River, floodplains, biogeochemical cycles, sediment dynamics, diagenesis, CO₂ outgassing.

1. Introduction

This paper is dedicated to the study of coupled biogeochemical, sedimentary and hydrological budgets of the Amazonian floodplains along a 2000 km-reach extending between the stations of Vargem Grande and Óbidos.

a. Preliminary work

Based on the chemical data of the CAMREX project (Carbon in the AMazon River Experiment), biogeochemical mass balances over the studied reach were calculated at 10 sampling sites, well spatially distributed, by comparing incoming and outgoing signals and fluxes (Bustillo, 2007) with respect to 44 physico-chemical parameters. This approach, based essentially on empirical observations instead of modeling outputs, emphasized that the anomalies of mass balances were mainly related to hydrograph stages and to the hydrological balance of the floodplains. Geochemical and hydrological information were treated in a lumped way, providing thus a pertinent insight on the complex hydrological and chemical linkages normally present between floodplains and river channels. Deliberately based on facts instead of modeling outcomes, this preliminary work raised many intriguing questions with respect to the structure of flux and signal anomalies (e.g. the coarse fraction of particulate organic carbon is very significantly ¹³C enriched during falling waters). The calculations of mass balances were performed in an

exhaustive way, at the 10 monitoring stations, involving 44 physico-chemical parameters over 8 sampling cruises. However, the determinants of flux imbalances remain to be identified.

b. Research objectives and challenges

This study aims at deciphering the nature and magnitude of underlying processes driving the transport of particulate and dissolved species towards the ocean and the gaseous exchanges at the river-atmosphere interface, with a special focus on floodplain-channel linkages. This last component is especially significant in large river systems and is highly relevant to contemporary international debates on the human modification of floodplain land-use, flood control, and the construction of levees and reservoirs which all act to de-couple floodplains from stream channel environments. Spatially linking tributary streams and longitudinal shifts in hydrology, water chemistry, and sedimentology is therefore a very challenging issue. Achieving this purpose requires testing the validity of the interpretative statements inspired by the empirical observations (preliminary work). The question addressed in this paper can be formulated as follows: what is the actual impact of the floodplains and that of their hydrological functioning on the biogeochemical budget of the Amazon River basin ?

To these ends, six hydrochemical modelling strategies, based on end-member mixing concept and using tracer-based separation methods, were implemented. These approaches aim to link hydrological pathways and chemical signals in order to couple hydrological and biogeochemical budgets. Comparing the outputs of these different models designed for different purposes is expected to provide a better sense of the whole by constraining better the range of possible interpretations given to the flux imbalances.

c. On the use of end-member mixing models

End-member mixing models provide comprehensive understanding of runoff generation processes with a special focus on hydrological pathways, contributive areas and retention times (Gonzales *et al.*, 2009). However, the direct measurement of each contributive runoff in a continuous way and at a sufficient number of locations is practically impossible (Tardy *et al.*, 2004). Hydrograph separation methods can be divided in two main categories: tracer-based and non-tracer-based separation methods. Non-tracer-based separation methods are based on the analysis of hydrographs, including a large variety of procedures among which: graphical analysis of recession curves, low pass filtering, unit hydrograph model with extrapolation to rising limb of hydrographs, and rating curve methods linking groundwater levels and river flow. Tracer-based separation methods are based on a mass balance approach

determined by the conservative mixture in variable proportions of compositionally constant end-members (or at least sufficiently stable and distinct from one end-member to another to make achievable the procedure). They are usually recognized to deliver valuable information about the groundwater contribution to the river discharge, provided that adequate tracers are selected. The procedure proposed by Hooper *et al.* (1990), called end member mixing analysis and based on the identification of end-members by principal component analysis (PCA), was widely applied for studying the hydrology of small catchments. Its implementation supposes that the water chemistry within each hydrological component is known. At the scale of very large river basins, no such input data can be measured, more particularly because of spatial heterogeneities of tracer concentrations and because the fluxes supplied by hydrological reservoirs are not systematically conservative due to in-stream processes and fluvial filtering in river corridors and floodplains (Meybeck and Vörösmarty, 2005).

The synthesis of Mortatti (1995) attempted to provide new insights into Amazon River hydrology by gathering biogeochemical and hydrological approaches. However, the hydrograph separation is based on two reservoirs only, and although very interesting, it proved to be insufficient, more particularly because it did not allow to capture the very significant influence of the floodplains on the biogeochemical budget of the Amazon River basin. Despite peculiar cases which are easily corrected case by case, it appears that most of large river basins, whatever their morphology and hydroclimatology, are reasonably modelled using the hydrograph separation concept, dividing the total discharge into - at least - 3 reservoirs.

d. Organization of the manuscript

In order to overcome the difficulties mentioned above, we proposed to investigate the hydrologic function of the Amazon River floodplains, between Vargem Grande (before the confluence of Rio Iça) and Óbidos (the outlet of the studied area), by applying six complementary modelling approaches (including end-member mixing models) to the successive sampling stations located along the main stem of the Amazon. These are based on: (1) variable regional sources with (model M2) and without (model M1) correction of inputs by small tributaries, (2) variable hydrological sources with 3 end-members (model M3) to determine their individual compositional evolution, with contrasted response depending on hydrograph stage, throughout their course in the floodplains, (3) variable hydrological sources with 3 end-members including a correction on the baseflow to account for in-stream biogenic transformations (model M4), and (4) mixed approaches (models M5 and M6) combining the regional variability of

chemical signals (between river basins) and the variability related to hydrological source (between contributing runoffs or end-members), taking into consideration the defaults of floodplains water balance.

The compositional changes of the chemical baseline in each individual reservoir, set in evidence by comparing their composition within incoming (tributaries) and outgoing (Amazon River reach) runoffs are more particularly analyzed. By determining hydrological sources and the magnitude of their individual compositional changes, this approach delivers a valuable and original insight on the main factors (hydrological source, water budget of the floodplains, nature of hydrobiological pattern i.e. photosynthesis vs. mineralization, air-water gaseous exchanges, etc.) driving the biogeochemical and sedimentary budgets of Amazonian floodplains.

2. Study area and dataset

The main physiographic structural elements of the basin include (i) the Precambrian, highly weathered Guyana and Brazilian shields, (ii) the Andean mountains to the west, (iii) the Andean alluvial foreland, and (iv) a large alluvial plain along the Amazon main stem. Soils in the lowlands are generally deep and highly weathered, with widespread covers of sandy podzols in the shields. The soils in floodplains (and alluvial regions around mainstems draining the Andes) are much less weathered, due to the continuous input of fresh sediments, delivered by physical erosion. The Amazon River and most of its large tributaries have developed extensive floodplains which are integral parts of the river systems (Richey *et al.*, 1997). After leaving the Andean foothills, the tributaries of the Amazon converge into a large sedimentary plain where they deposit large volumes of sediments (*e.g.* Guyot, 1993) and inundate the floodplains via an extensive network of drainage channels called *paranas*. Two distinct types of floodplain channels are (i) tributary channels that drain upland terraces and (ii) distributary channels that transport main-stem Amazon water and sediment to floodplain lake basins. A synthetic view of the Amazon River basin upstream from Óbidos, including the delineation of major sub-basins, the location of sampling station along the Amazon River main stem and the geographical repartition of small tributaries ungauged during the CAMREX project, is presented (**Figure 1**).

The samples were collected during the CAMREX Project (Carbon in the Amazon River Experiment), over the period 1982-1984 (8 cruises), during contrasted hydrographic stages, completed by 5 additional cruises between 1985 and 1991 focusing on specific topics, *i.e.* for which exhaustive data set is not available (thus not considered in this paper). The objective of CAMREX project was to define by mass balances and direct measurements those processes which control the distribution of bioactive elements (C, N, P and O) in the mainstem of the Amazon

River in Brazil. The CAMREX dataset represents a time series unique in its length and detail for very large river systems. The data set, extracted from Pre-LBA compilation (Marengo and Victoria, 1998, available on-line: Richey *et al.*, 2008), consists in representative flux-weighted water samples for comprehensive chemical analysis measured over 18 different sites within a 2000 km reach of the Brazilian Amazon mainstem, including 7 major tributaries. This data set constitutes, until that date, the basis of more than 130 CAMREX publications which have focused on understanding physical and biogeochemical dynamics throughout the basin using a large variety of approaches (e.g. Richey *et al.*, 1990). Monitoring stations are located a few kilometres upstream of the confluence of the 7 major tributaries with the Amazon River: Iça, Japura, Jutai, Jurua, Purus, Negro and Madeira, and (2) along the Amazon River at the 11 following stations: Vargem Grande, Santo Antonio do Iça, Xibeco, Tupe, Jutica, Anori, Itapeua, Manacapuru, São Jose da Amajari, Paura and Óbidos (4 619 000 km²), the outlet of the studied area. Thus, it gets possible to compare the inputs from tributaries and the outputs of the Amazon River at different locations along the longitudinal profile of the mainstem.

3. Modelling strategy

Six mixing-models of increasing complexity are implemented (**Table 1**). They belong to three distinct categories. The first category (models M1 and M2) accounts the variable contribution of the sub-basins to the biogeochemical budget. The second category relies on end-member mixing models (models M3 and M4) which allow the identification of source reservoirs, supposed to have constant composition but contributing in variable proportion to the river flow. A third category of model, taking into account the variability related to regional contrasts and hydrological source, is also explored (models M5 and M6).

a. Variable regional source

The 1st model (M1) fundamentally relies on the comparison between calculated and observed longitudinal profiles of concentrations. Upstream, the chemical composition of Rio Solimões constitutes the starting point (1), and along its course in low plains, several tributaries join the principal river and modify its chemical composition: Rios Iça (2), Jutai (3), Japurá (4), Juruá (5), Purus (6), Negro (7), Madeira (8), the last one before Óbidos, the outlet of the Amazon River basin (9) chosen in this study. Considering the concentrations (C_i^j) of each chemical species i , the

total discharge (Q_t^j) of the j^{th} confluent, and before the confluence with the tributary ($j + 1$), the concentration $C_{i,tot}^{j+1}$ in the Amazon River after the $j+1^{\text{th}}$ confluent, is established as follows:

$$C_{i,tot}^{j+1} = \left[C_{i,tot}^j \times \sum_{k=1}^{k=j} Q_t^k + C_i^{j+1} \times Q_t^{j+1} \right] / \sum_{k=1}^{k=j+1} Q_t^k \quad (1)$$

Pedices i and apexes j correspond respectively to the parameter analysed and to the number of tributaries contributing to the Amazon River flow at each station considered (from $j = 1$ standing for Santo Antonio do Içá to $j = 8$ designating the station of Óbidos). Assuming conventionally that $Q_{tot}^j = \sum_{n=1}^{n=j} Q_{tot}^n$ and $Q_{tot}^{j+1} = \sum_{n=1}^{n=j+1} Q_{tot}^n$, one obtains:

$$C_{i,tot}^{j+1} = \left[C_{i,tot}^j \times Q_{tot}^j + C_i^{j+1} \times Q_t^{j+1} \right] / Q_{tot}^{j+1} \quad (2)$$

After the j^{th} confluent, $C_{i,tot}^{j+1}$ becomes $C_{i,tot}^j$, and the cumulative runoff changes from Q_{tot}^{j+1} to Q_{tot}^j .

For each parameter, the concordance between theoretical (i.e. calculated) and observed longitudinal profiles is appreciated by analysing the fitting capability of the simple linear model:

$$C_i^j (obs.) = \alpha_i^j \times C_i^j (calc.) + \beta_i^j \quad (3)$$

The correlation coefficient R^2 is calculated for each parameter ($n_i = 44$) and each station ($n_j = 10$). The results of calibration are given in appendix (Table A2). Then, the slope α_i^j (ideally close to 1) and the intercept to origin β_i^j (ideally close to 0) are considered. Finally, the mean bias B is calculated as follows:

$$B_i^j = \frac{\sum_{k=1}^{k=Nk} C_{i,k}^j (calc.) - \sum_{k=1}^{k=Nk} C_{i,k}^j (obs.)}{N_{i,k}^j} \quad (4)$$

where k indexes the number of the sample and $N_{i,k}^j$ for the number of samples for each station and each parameter. This bias, calculated on average discharge-weighted values, allows estimating the mean chemical composition of floodplains and small tributaries $C_i^j(v)$, assuming that measured biases depend on their variable contribution to river flow. The model M2 corrects M1 by taking into account the mean composition of small

tributaries and floodplains. The composition of the additional flow (small tributaries + alluvial aquifers), noted $C_i^j(v)$, is estimated by the analysis of differences (composition and flow) between the sum of major tributaries

(calculated input): $\Phi_i^j(in) = C_i^j(in) \times Q_i^j(in)$ and the output (measured output):

$\Phi_i^j(out) = C_i^j(out) \times Q_i^j(out)$ corresponding to the river water composition at the considered station.

$$C_i^j(v) = (\Phi_i^j(out) - \Phi_i^j(in)) / Q_i^j(v) = (C_i^j(out) \times Q_i^j(out) - C_i^j(in) \times Q_i^j(in)) / (Q_i^j(out) - Q_i^j(in)) \quad (5)$$

Finally, the concentrations obtained through the model M2 for each cruise are given by:

$$C_{i,k}^j(M_2) = \{ C_{i,k}^j(M_1) \times Q_{t,k}^j + C_i^j(v) \times [Q_{t,k}^{j+1} - Q_{i,k}^j] \} / Q_{t,k}^{j+1} \quad (6)$$

In order to test the model fitting capability, linear equations comparable to those presented above for M1 are also calibrated for M2. By averaging the estimations of all the samples, it must be reminded that $B_i^j = 0$ for each station and each parameter, in the case of M2.

b. Variable hydrological source.

The third model (M3) relies on the hydrograph separation into three components: *RS*, the superficial runoff, *RI* the interflow and *RB* the baseflow. These reservoirs are meant as the expression of spatially organized tributary basins with vertical (top-bottom in soils) and upstream-downstream gradients involving three mixing end-members:

- *RS* tracking superficial and hypodermic pathways that arise more particularly in upstream areas which provide most of the solid load transported by fluvial systems;
- *RI* tracking superficial and hypodermic pathways that arise more particularly in downstream areas where deep leached soils provide waters of low dissolved (except for organic matter) and solid loads;
- *RB* tracking groundwater pathways, corresponding to the leaching of the soil horizon C (permeable saprolite) that occasionally emerge in the gleysols of lowlands, as inferred from the very characteristic ^{18}O enrichment of waters (Tardy *et al.*, 2009).

The identification of these three components relies upon chemical tracing. The concentrations of Na^+ and FSS (fine suspended sediment) are selected as the best tracers (Tardy *et al.*, 2005). The fluctuations of $[\text{Na}^+]$ track the processes of dissolution and evaporation which tend to generate a concentration gradient from the superficial layer of soil to groundwaters which are directly at the contact with the chemical front of alteration. On the other side, the

fluctuations of [FSS] track the soil erosion which is almost specific of surface runoff. The chemical tracers in each reservoir determine the contribution of these source reservoirs to the total river flow Q_t by solving the system of equation composed of 2 equations of mass conservation for each tracer (Eq.7) and the equation of flow conservation (Eq.8).

$$C_{i,k}^j = C_i^j(Rs) \times QRs_k^j / Q_t^j + C_i^j(Ri) \times QRi_k^j / Q_t^j + C_i^j(Rb) \times QRb_k^j / Q_t^j \quad (7)$$

$$QRs_k^j + QRi_k^j + QRb_k^j = Q_t^j \quad (8)$$

where j indexes the considered sub-basin and i the parameter used as a chemical tracer.

[Na⁺] and [FSS] within each reservoir corresponds to the values established by Tardy *et al.* (2005). The next step consists in adjusting $C_i^j Rs$, $C_i^j Ri$ and $C_i^j Rb$ to the whole data set (42 parameters excluding the 2 tracers) by performing multilinear regressions. As a result, we define statistically the most probable composition of each reservoir Rs , Ri , Rb . End-member mixing models are calibrated for each of the 10 stations located along the Amazon River main stem (from Saõ Antonio do Içá to Óbidos) but also for each of the 8 major tributaries. Then, the decomposition of hydrograph is performed theoretically by cumulating the reservoir inflow Q_K of each tributary (*trib*) to each station ($1 \leq j \leq 8$).

$$Q_K^{j+1} = Q_K^j + Q_{K,trib}^{j+1} \quad (9)$$

where κ is the index for hydrological reservoirs : Rs , Ri or Rb . Consequently, for each hydrological node, two models of repartition are implemented. The first one is calibrated using chemical data measured at each of the 10 stations, whereas the second one is calibrated using a calculated pool of chemical data, corresponding to the variable spatial contribution of sub-basins to the Amazon River discharge. It is expected that differences between the chemical characteristics of reservoirs are good indicators of floodplains biogeodynamics. Full data is provided in appendix (Table A3).

c. Biologically-mediated processes

The biological control of chemical factors in river, popularized by Redfield (1958), is evaluated in the model M4 by testing the influence of biotic processes on the composition of the baseflow Rb . The protocol of calculation for evaluating $QRs(j)$, $QRi(j)$ and $QRb(j)$ is identical to that of the model M3. The composition of baseflow is supposed to be variable as a function of biological pathway tracked with the synthetic variable $IBIO$:

$$I_{BIO\ k}^j = [O_2]_k^j - [CO_2]_k^j \quad (10)$$

In the case of intense photosynthesis, O_2 is actively produced while CO_2 is removed and consequently I_{BIO} increases. Conversely, when the decomposition prevails, CO_2 is actively produced while O_2 is removed and consequently I_{BIO} diminishes. The model M4 is formalised as follows:

$$C_{i,k}^j = C_i^j(Rs) \times QRS_k^j / Qt_k^j + C_i^j(Ri) \times QRI_k^j / Qt_k^j + [C_i^j(Rb) + K_{BIO\ i}^j \times I_{BIO\ k}^j] \times QRB_k^j / Qt_k^j \quad (11)$$

with $K_{BIO\ i}^j$ corresponding to the rate of uptake or release of each bioactive element (i) for each station (j) associated to biologically-mediated processes in the river water. If $K_{BIO\ i}^j > 0$, the concentration increases when the photosynthesis pathway prevails ($I_{BIO\ k}^j > 0$) and decreases when the mineralization predominates ($I_{BIO\ k}^j < 0$). The mineralization leads to the removal of dissolved O_2 and to the release of CO_2 . That is the reason why I_{BIO} associated to mineralization paths is usually negative and potentially very negative. Therefore, $K_{BIO\ i}^j < 0$ indicates that the concentration increases when mineralization pathway prevails ($I_{BIO\ k}^j < 0$) and decreases when photosynthesis predominates ($I_{BIO\ k}^j > 0$). In the case of isotopic data ($\delta^{18}O$, $\delta^{13}C$) which are all negative, the interpretation of $K_{BIO\ i}^j$ is inverted. For simplification purposes, the signs of $K_{BIO\ i}^j$ associated to isotopic values were systematically inverted to homogenise the deciphering for all the parameters. Full model outcomes relative to M4 are presented in appendix (Table A4).

d. Composite approach

The model M5 is a composite approach which integrates both variable spatial contribution (M1) and variable hydrological source (end-member mixing models). First, we establish for each cruise k ($Nk = 8$), each parameter i and each hydrological node j , the relative difference, noted $\Delta C_{i,k}^j$, between calculated (calc.) and observed (obs.) values, as follows:

$$\Delta C_{i,k}^j = \frac{C_{i,k}^j(obs.) - C_{i,k}^j(calc.)}{C_{i,k}^j(calc.)} \quad (12)$$

The second step consists in relating this relative difference with several factors. We have selected three covariates corresponding to the relative differences $\Delta(QRS / Qt)$, $\Delta(QRI / Qt)$ and ΔQt :

$$\Delta C_{i,k}^j = \alpha_i^j \times \Delta(QR_s / Qt)_k^j + \beta_i^j \times \Delta(QR_l / Qt)_k^j + \gamma_i^j \times \Delta Qt_k^j + \delta_i^j \quad (13)$$

$$\text{with: } \begin{cases} \Delta(QR_s / Qt)_k^j = (QR_s / Qt)_k^j, M_3 / (QR_s / Qt)_k^j, M_1 - 1 \\ \Delta(QR_l / Qt)_k^j = (QR_l / Qt)_k^j, M_3 / (QR_l / Qt)_k^j, M_1 - 1 \\ \Delta Qt_k^j = Qt_k^j, \text{obs} / Qt_{tot,k}^j - 1 \end{cases} \quad (14)$$

δ_i^j in Eq.13 stands for the residual relative difference $\Delta C_{i,k}^j$ when the three following conditions are fulfilled:

$$(i) : \Delta(QR_s / Qt)_k^j = 0 ; (ii) : \Delta(QR_l / Qt)_k^j = 0 ; (iii) : \Delta Qt_k^j = 0$$

The coefficients α_i^j , β_i^j and γ_i^j , estimated by multilinear regressions, provide qualitative information on river diagenesis in Rs (surface runoff), Rl (interflow) and Qt (total runoff). If the coefficient is positive, it indicates that the concentration increases in the correspondent runoff as the individual discharge Q_k increases. Considering the total river flow Qt , the sign of γ_i^j indicates whether the discharge of floodplains, roughly estimated by $\Delta Qt_k^j = Qt_k^j / Qt_{tot,k}^j - 1$ contributes to increase or decrease the chemical concentration of the chemical parameter i in the river water at the station j . Values of QR_s, QR_l, QR_B and Qt are given in appendix (Table A1 and Figure A1), M3 corresponding to model-derived data and M1 to data calculated from upstream sub-basins. A complementary approach (M6) consists in evaluating the combined effect of the total discharge Qt_k^j and its excess or deficit ΔQt_k^j :

$$\Delta C_{i,k}^j = \alpha_2^j \times \Delta Qt_k^j + \beta_2^j \times Qt_k^j + \gamma_2^j \times Qt_k^j \times \Delta Qt_k^j + \delta_2^j \quad (15)$$

The calibration of these 4 coefficients (α_2^j , β_2^j , γ_2^j and δ_2^j) for each sampling station and each parameter leads to synthetic 3-D diagrams (see Figure A2 in appendix) which allow describing the compositional fluctuations of the river water as a function of the river discharge Qt_k^j and the default of water balance ΔQt_k^j .

4. Results and Discussion

After a brief comparison of the 5 models in term of statistical resolution capability, the information supplied individually by each method is analyzed. The analysis of their mutual consistency is more particularly performed.

a. Compared performance

The agreement between calculated and observed water composition is very significant for most of the parameters. The cumulative distribution of correlation coefficient for 5 tested models is presented **Figure 2**. This indicates that 50% of parameters modelled by M1 exhibit $r^2 > 0.75$. Unexpectedly, M1 provides better results than M2 does, suggesting that floodplains are not of constant composition, contrary to the assumption underlying the approach M2. The comparison between M3 and M4 reveals a significant improvement of the performance of end-member mixing models by taking into account the “hydrobiological index” I_{BIO} which allows identifying the parameters influenced by in-stream processes. The level of performance remains deficient (threshold arbitrarily fixed to $r^2 < 0.60$) on 25 % of parameters for M4 vs. 50% of parameters for M3. Finally, the mixed approach (M5) combining variable spatial contribution and end-member mixing models constitutes a very convenient compromise which provides the best results.

b. Variable regional contribution (M1 and M2)

The simple approach consisting in correlating incoming (theoretical, calculated) and outgoing (measured) concentrations (M1) provides insightful information. It appears that most of linear calibrations are very significant, except for SO_4^{2-} , HPO_4^{2-} , CSS, POCC, C/N (in all fractions) and PONC. These deficiencies reveal that substantial modifications occur in the floodplains.

Table 2 delivers the mean values of α (slope), β (intersect of line for $x = 0$), r^2 (correlation coefficient), bias and average for each parameter, and for the 10 sampling stations located along the Amazon main stem. These linear calibrations indicate that the compositional fluctuations of river water in the Amazon reach might be greater than those impulsed by the tributaries inputs ($\alpha > 1$ and $\beta < 0$) for pH, K^+ , Mg^{2+} , NO_3^- , CO_2 , NaSil, KSil, CaSil, MgSil, dolomite and FR . For example, when the inflow defines a low pH, the outflow is still more acidic and conversely when the inflow defines an elevated pH, the outflow is more basic. Considering the parameters listed before, the open system dynamics along the Amazon main stem (and in its floodplains) accentuate the chemical perturbations initiated upstream, in the sub-basins. Conversely, the compositional fluctuations generated in the tributaries tend to be buffered in the outflow ($\alpha < 1$ and $\beta > 0$) for other parameters such as Ca^{2+} , HCO_3^- , DIC, Cl^- , DOC, O_2 , CSS, POCC, C/N (all the fractions), PONF, PONC, DON and $[\text{CaCO}_3]$.

A complementary analysis of **Table 2** consists in assessing the bias between the concentrations in the inflow and in the outflow. Positive values (Bias > 0) indicate that concentrations in the inflow are superior to those measured in the outflow, and vice versa. Negative biases are observed for Ca^{2+} , HCO_3^- , Cl^- , O_2 , CSS, PONF, PONC and

CaCO₃. We observe also a decrease of $\delta^{13}\text{C}$ in all the carbon fractions: DIC, POCF and to a lower extent POCC. In turn, positive biases are obtained for NO₃⁻, CO₂ and C/N (in all the fractions) while $\delta^{18}\text{O}$ gets less negative. Globally speaking, the measured composition of the Amazon river (outputs) follows the chemical baseline imprinted by the tributaries (inputs). Thus, in-stream processes arising in the studied reach do not modify fundamentally the chemical composition acquired in the tributaries. The mitigation of compositional fluctuations is probably related to the contribution of ungauged rivers which influence substantially the chemical signal measured in the Amazon River (e.g. Ca²⁺, HCO₃⁻ and Cl⁻) due to the very low salinity of small rivers draining thick-sandy soils in Central Amazonia. Conversely, the accentuation of trends observed downstream seems to be due to organic matter decay which is expected to take place in the floodplain as water slowly enters the stream-channel from temporary storage. This leads to the release of CO₂ (¹³C depleted) and nitrogenous dissolved species (NO₃⁻, DON), and symmetrically to the removal of O₂.

In the model M2, outputs are adjusted by prescribing ad-hoc additional contribution of small rivers (whose average composition is not accurately known) which border the Amazon River. The reconstituted mean annual composition of small rivers and floodplains (Bustillo, 2007) delivers reliable results for most of the parameters and provides valuable insight on the presumed impact of river processes. Yet, the correction proposed in the model M2, relying on the variable contribution but constant composition of small rivers and adjacent floodplains, unambiguously fails (**Figure 2**). This may be due to several factors: (1) the area of flooded areas, on which fall direct precipitation, is variable, (2) the biotic transformations undergone by transiting materials are not the same depending if floodplains fill or dry up, and (3) the respective contributions of small rivers draining lowlands (low TDS) and groundwater (high TDS) fluctuate along the hydrological cycle.

c. Hydrograph separation into 3 reservoirs (Model M3)

The detailed modelling outcomes relative to M3 are presented in appendix (Table A3). A simple comparison can be made between inputs (calculated) and outputs (observed). Mean values are more accurately examined (**Figure 3a** and **Figure 3b**). Observed values are established by averaging the results obtained for 9 sampling stations (excluding Vargem Grande and Santo Antonio do Iça which are located at – and close to - the upstream boundary) located along the Amazon River main stem. Calculated concentrations ($\hat{C}_{i,k}$) are obtained as follows:

$$\hat{C}_{i,k} = 1/9 \times [C_{i,k}^{j=2} + C_{i,k}^{j=3} + 3 \times C_{i,k}^{j=5} + C_{i,k}^{j=6} + C_{i,k}^{j=7} + 2 \times C_{i,k}^{j=8}] \quad (16)$$

Notice that: $[C]R_S \times QR_S/Q_t + [C]R_I \times QR_I/Q_t + [C]R_B \times QR_B/Q_t = [C]_{AVE}$ for all the chemical parameters.

The agreement between both data sets is good, except for C/N, DON, HPO_4^{2-} , CO_2 , O_2 and pH. Despite some unavoidable deviations due to the imprecisions of the chemical analyses and due to simplifying assumptions required for modeling, the repartition of chemical species and isotopic signatures display the same pattern. The compositional contrasts between the 3 reservoirs tend to decrease in the outflow, suggesting that the intermittent storage of water in floodplains contributes to mix waters originating from different sources (hydrological reservoirs, sub-basins). This tends to homogenize their chemical composition at the outlet of the system. The examples provided by CSS, $\delta^{18}O$, DOC (**Figure 3b**), SO_4^{2-} and Cl^- (**Figure 3a**) are particularly explicit. The greatest deviations are observed for the sand fraction CSS whose transport in the Amazonian reach is considerably delayed compared to solutes and water.

Concerning the major chemical species (anions and cations), concentrations in R_S and R_I tend to be lower in the outflow: this effect of dilution is very marked for SO_4^{2-} , Cl^- , DOC, Ca^{2+} and HCO_3^- (**Figure 3a**). This is consistent with the biogeochemical balance calculated for the floodplains (Bustillo, 2007) which did not reveal any dissolution of carbonates in Central Amazonia. It is likely that a part of Ca^{2+} is adsorbed on transiting clay suspensions, fulvic acids and/or goethite (Weng *et al.*, 2005) while HCO_3^- might be partly converted into CO_2 as a result of pH buffering of very acidic waters provided by small Amazonian rivers draining lowlands and Negro river. We observe also a very significant increase of weathering rate in the outflow (see CO_2 SIL), attributed to the baseflow R_B . The consequence is the correlative decrease of the lithological index FR and $\delta^{13}C$ (DIC): $-9.6\text{‰} \rightarrow -11.7\text{‰}$. However, the values established for $\delta^{13}C$ (DIC) and FR are not totally compatible because a low contribution of carbonates on DIC release (FR low) should lead to a very negative $\delta^{13}C$ (DIC). The unexpected heavy signature of $\delta^{13}C$ (DIC) in the baseflow (**Figure 3b**) might be the consequence of (1) CO_2 outgassing (Richey *et al.*, 2002) and (2) aquatic photosynthesis, which both subtract preferentially ^{12}C and thus concentrate ^{13}C in the river water.

Focus on the sampling stations of Paurá and Óbidos

The outcomes of end-member mixing model are more particularly analyzed downstream from the confluence of the 8 major tributaries, at the monitoring stations of Paurá and Óbidos (**Table 3**). Chemical characteristics of the three reservoirs are almost similar for Paurá and Óbidos; they display distributions comparable to those resulting from the conservative mixing of the 8 major tributaries. Even so, the statistical resolution might be deeply altered on

specific parameters for one station and not for the other (*e.g.* Ca^{2+} , Mg^{2+} and HCO_3^-). Small errors of chemical analyses might have large repercussions on the outcomes of the models, especially when the number of samples is low and when the variability of the chemical baseline is moderate. For the sampling stations located at the outlet of large fluvial basins, the compositional fluctuations of the river water are often attenuated, due to the water storage in the floodplains and the slow motion of floodwave which mixes waters having resided for short or long time in the surface network. Even when r^2 are low, the model outcomes are qualitatively very instructive to appreciate the dynamics within each reservoir and their heterogeneity. The similarity between the fictitious station (8 Rios) and the sampling stations (Paurá and Óbidos) indicates that the chemical composition (including isotopic composition) of the Amazon River water is essentially acquired before the waters supplied by tributaries reach the Amazonian floodplains. It seems that underlying processes driving biogeochemical budgets (chemical weathering, gas emissions towards the atmosphere, deposition vs. remobilization of sediments, etc) in the tributaries and in the Amazon main reach are of same nature and define comparable chemical equilibria. The large-scale flooding of lowlands, occurring almost concomitantly over Central Amazonia (Hamilton *et al.*, 2002), provides autochthonous organic substrate for decomposition, leading subsequently to the creation of hypoxic and anoxic environments along river corridors. These reductive conditions influence considerably carbon and nutrient cycling due to the enhancement of gas emission (CH_4 , CO_2 , NO , N_2O , N_2) towards the atmosphere and thus determine the isotopic signature of the dissolved inorganic carbon. The major chemical species released by chemical weathering in upstream reaches, where the bedrock outcrops, and in the floodplains, where coarse unweathered sediments deposit, seem to be transported (almost) conservatively within the river network.

d. Hydrobiological modelling (Model M4)

The analysis focuses more particularly on the sign of the coefficient $K_{BIO\ i}^j$ associated to the hydrobiological factor $I_{BIO\ k}^j$. Model outcomes, on several characteristic parameters and for the 11 monitoring stations of the Amazon profile are represented at **Figure 4** (full dataset at Table A4, electronic supplementary material). Chemical responses to hydrobiological factor may be roughly grouped into three categories: (1) $K_{BIO\ i}^j > 0$, indicating that concentrations (or values) increase concomitantly to I_{BIO} , or that concentrations are higher when photosynthetical paths dominate; (2) $K_{BIO\ i}^j < 0$, indicating that concentrations (or values) increase when the mineralization prevails on the photosynthesis; (3) $K_{BIO\ i}^j \ll \overline{C_i^j}$, indicating that hydrobiological processes do not influence

significantly the chemical baseline. Among the parameters varying like I_{BIO}^j , we have: pH, K^+ , NO_3^- , SO_4^{2-} , HPO_4^{2-} , O_2 , $\delta^{13}C$ (DIC), $\delta^{13}C$ (POCF) and C/N (3 fractions). Among the parameters varying in the opposite sense, we have: Ca^{2+} , HCO_3^- , DOC, DIC, CO_2 , POCF, PONF, PONC, DON. Other parameters do not exhibit any significant and reproducible correlation with I_{BIO}^j . A break is observed at Manaus, depending if we locate upstream or downstream from the confluent of Rio Negro. Roughly speaking, there is a loss of Ca^{2+} , Mg^{2+} , K^+ , HCO_3^- under photosynthetic regime and conversely a gain of these solutes after the confluent of Rio Negro. Similarly, the decrease of [DOC] is clearly observed downstream from Manaus but remains negligible upstream.

When the photosynthetical paths dominate ($I_{BIO}^j \gg 0$), CO_2 is removed while O_2 is released. As a result, the pH increases and influences the nature and magnitude of abiotic processes. First of all, the biological uptake of CO_2 operates a carbon fractionation which tends to make heavier, by mass effect, the value of $\delta^{13}C$ (DIC) in the river water. Concomitantly, the decrease of $[Ca^{2+}]$ and $[HCO_3^-]$ supports the hypothesis that Ca^{2+} is adsorbed by clay minerals, which releases H^+ and leads to the subsequent protonization of HCO_3^- (which frees CO_2). After the confluence of Rio Madeira, the dynamics of Ca^{2+} is reversed: all other things being equal, $[Ca^{2+}]$ and $[HCO_3^-]$ tend to increase (photosynthetical path). The rise of pH and/or $[O_2]$ related to photosynthetical paths might promote the side-chain oxidation of nitrogenous functions contained in dissolved organic molecules (Aufdenkampe *et al.*, 2001; Aufdenkampe, 2002). Their condensation makes them get more hydrophobic (Tardy *et al.*, 2009), and leads presumably to their sorption onto fine suspended sediments to form diagenetic POCF. The rise of POCF/PONF and POCC/PONC reflects the genesis of autochthonous molecules which progressively obliterates the signal of soil-derived substances. Actually, riparian grasses and floating aquatic plants which grow in the floodplains exhibit high atomic ratio C/N, evaluated to 42 by Victoria *et al.* (1992), suggesting that the uptake of NO_3^- is rather low and does not counter-balance totally the input associated to the sorption of DOM. However, the most significant rise is observed for [DOC]/[DON] due to the diagenesis of DOM which tends to release NO_3^- and to concentrate carbon in DOM.

Conversely, when the mineralization paths prevail on photosynthesis ($K_{BIO}^j \ll 0$), low $^{13}C/^{12}C$ source of DIC is released in the river while O_2 is consumed. The chemical signals associated to mineralization are, roughly speaking, symmetrical to those imprinted by the photosynthesis. The Andean soil-derived POCC, mainly mobilized in surface runoff, is exposed to increasing temperature as transported downstream and subjected to mineralization in lowland environments (McClain *et al.*, 1995). This source of unstable carbon provides significant amount of

carbon substrate that fuels the heterotrophic metabolism of the river, after deposition of large quantities of carbon. In addition to allochthonous POCC, the mechanism locally named “terras-caídas” corresponding to the large-scale bank erosion during flood periods (Irion *et al.*, 1997), promotes the large-scale destruction of well developed floodplain forest communities (Junk and Piedade, 1997) and provides large amounts of highly unstable organic substrate. The contribution of grass várzeas to the carbon budget of floodplains, appreciated by the isotopic composition of várzea sediments (Victoria *et al.*, 1992; Martinelli *et al.*, 2003), increases as we move downstream. This suggests that the impact of aquatic vegetation on the carbon budget of floodplain progressively increases (Quay *et al.*, 1992). Owing to preservation mechanisms during decomposition (such as adsorption-linked protection), the fine fraction POCF is clearly refractory: $K_{BIO}^j << [POCF]_{AVE}$, and does not appear to be significantly influenced by in-stream mineralization.

e. The composite approaches (M5 and M6)

The model M5 exhibits an unexpected high performance, for all monitoring stations and all chemical parameters, as shown in appendix (Table A5). The reconstitution of isotopic signatures is very convincing for $\delta^{18}O$ (SMOW) and $\delta^{13}C$ (DIC) while the lowest levels of confidence are observed for POCC, CSS, PONC and C/N atomic ratios for POCC and POCF. The significant improvement compared to M1 and M3 denotes the influence of floodplains on the chronological variations of the Amazon River composition, mainly due to (1) the polarity of water circulation in the floodplains, and (2) to the variable contribution of each individual runoff to the water budget of floodplains. The chemical response of river to the polarity of water circulation in the floodplains can be approached by analysing the magnitude and the sign of γ_i^j (cf. Eq.13) corresponding to the variation of concentration in the river water associated to the water balance of floodplains $\Delta Q t_k^j$. The complete data set compiling the values of γ_i^j is supplied in the electronic supplementary materials (Table A5); several selected values are represented at

Figure 5. Despite some variations between stations, the magnitude and the sign of γ_i^j match quite well. The model M5, taking into account simultaneously the following variables: $\Delta Q t_k^j$, $\Delta QRS/Q t$ and $\Delta QRt/Q t$, although providing reliable outcomes, might sometimes be difficult to interpret, notably because of covariations between variables. For example, $\Delta Q t_k^j$ and $\Delta QRt/Q t$ exhibit a positive correlation simply because the drainage of

floodplains involves principally the hydrological reservoir R_l . Moreover, the simple interpretation of γ_i^j does not allow investigating the additional impact of the river flow, whose magnitude influences floodplain dynamics.

To facilitate the deciphering process, the model M6 was implemented. The calibration of the four parameters (see Eq.15), leads to the determination of 3-D diagrams with spatial representation realized for $600 \text{ mm.yr}^{-1} \leq Qt_k^j \leq 1600 \text{ mm.yr}^{-1}$ (average = $1\,122 \text{ mm.yr}^{-1}$) and $-0.05 \leq \Delta Qt_k^j \leq 0.30$ (average = 0.134). These figures, available in full in appendix (Figure A2) for the station of Óbidos, represent the variations of concentration (or isotopic value) as a function of the river outflow, and as a function of the water balance of floodplains, tracked by ΔQt_k^j . The mean concentration of each parameter, calculated in the inflow (for 8 cruises), is centred on $Qt_k^j = 1\,122 \text{ mm.yr}^{-1}$ and $\Delta Qt_k^j = 0.134$. The isolines represent the changes of chemical characteristics as we stray from the central point, assuming that the chemical composition of the inflow remains constant whatever Qt_k^j and ΔQt_k^j . In this section, the objective is not to predict the chemical composition of the outflow but rather the deviation to the inflow. As a consequence, the following diagrams must be read and interpreted only in terms of relative values. On each diagram are identified the main sequence of hydrological cycle : (1) lowest waters, characterised by low Qt and excess of outflow \rightarrow (2) rising waters, exhibiting intermediate Qt and a severe deficit of outflow \rightarrow (3) highest waters, with high Qt and moderate deficit of outflow \rightarrow (4) falling waters, with intermediate Qt and excess of outflow \rightarrow (1) lowest waters.

1) MAJOR CHEMICAL SPECIES AND CHEMICAL WEATHERING

Concerning the major chemical species, a dilution effect is observed during the phase of water storage while concentrations rise when the waters stored in the floodplains join back the main channel. The influence of the circulation polarity is greater after the confluence of Rio Japurá, as the Amazon valley widens. Considering the poles of chemical erosion, it is noticeable that the apparent rate of chemical alteration is substantially greater (both for silicates and carbonates) when the floodplains empty, with the exception of dolomite, all other things being equal (**Figure 5**). On the average, the lithological index FR (cf. list of parameters) exhibits lower values during the emptying of floodplains. It suggests that submerged low plains drain areas where the alteration of sedimentary minerals mimics the weathering of crystalline rocks. It is likely that chemical weathering in floodplains manifests itself sequentially, when the floodplains dry up i.e. when the water stored in low plains joins back the main

channel. Following Johnsson & Meade (1990), these model outcomes support the idea that the chemical weathering of the additional flow is mainly driven by the diagenesis of unweathered sediments, deposited during the filling of floodplains (Martinelli *et al.*, 1993). The decrease of lithological index Fr during the emptying of floodplains coincides with lower $\delta^{13}C$ (DIC), compared to the stage of filling (**Figure 6a**). The parallel evolution of $\delta^{13}C$ (DIC) and Fr confirms that the isotopic signal of dissolved inorganic carbon is fundamentally determined by the pattern of weathering processes (synthesized by Fr). Because of the hydrological dynamics of floodplains, following an annual immutable cycle, the chemical expression of weathering processes is sequential as well. It should also be noticed that the index Re (SiO_2/Al_2O_3 in altered products inferred from the chemistry of river water, Tardy *et al.*, 2004) is much higher during the emptying of floodplains, suggesting that SiO_2 might be picked up and converted into a particulate form due to the growth of diatoms which is encouraged in adjacent lakes and flooded areas.

2) BUDGET OF SEDIMENTS

The sand fraction CSS exhibits systematically lower concentrations during the phase of water storage (γ_i^j are all positive at **Figure 5**) than during the emptying stage (**Figure 6b**). This result supports the idea that the sediments tend to deposit as the river inundates the low plains and tend to be remobilized as the extension of submerged areas lessens. The greatest contrasts are observed between Itapeua and São José da Amajari and the lowest is accredited to Paurá, after the confluence of Rio Madeira which provides large amounts of coarse sediments. The trends for the silt-clay fraction are less explicit. Considering the silt-clay fraction FSS, no clear tendency can be outlined on the Vargem Grande-Manacapuru reach. In turn, between São José da Amajari and Óbidos, the pattern is very similar to the one described for CSS, suggesting that the flushing action of Negro River (Meade *et al.*, 1985; Dunne *et al.*, 1998) might promote the remobilization of fine sediments when the floodplains drain.

3) RIVER METABOLISM

Considering the gaseous composition of the river, the computed γ_i^j indicate higher $[CO_2]$ during the emptying of flooded area on the upstream reach and lower $[CO_2]$ on the downstream reach (from Manacapuru to Óbidos). All other things being equal, pH appears to be higher during the emptying of floodplains while $[O_2]$ is lower.

Considering nitrogenous species, the evolutions are the following: while the floodplains dry up, a generalized drop of $[\text{NO}_3^-]$ and a gain of $[\text{PONC}]$ and $[\text{DON}]$ are observed, with the exception of Paurá and São José da Amatari where sorption processes of dissolved organic matter arise (Aufdenkampe *et al.*, 2001; Tardy *et al.*, 2009).

Considering organic carbon species, γ_i^j highlight a gain of POCF and POCC during the emptying of floodplains and conversely a deficit of DOC. The major exception corresponds to the station of Paurá, highly influenced by the forwarded contribution of Rio Madeira which reverberates directly on the chronological evolution of processes.

As the waters stored in the floodplains join back the main channel, the atomic ratio POCF/PONF rise whereas DOC/DON and POCC/PONC drop drastically. The drift observed along the Amazon profile involves the increasing contribution of submerged areas to the water budget of the Amazon River, as we move downstream. This amplifies the imprint of river diagenesis on the organic matter. As heterotrophic processes operate more and more intensely, the maturation of organic matter is accelerated and leads to lower and lower POCF/PONF in the water draining floodplains. The effects of these processes on the isotopic signature of carbon are not appreciable. Concerning POCC, the concentrations tend to be higher during the emptying of the floodplains. The effects on the isotopic signal of $\delta^{13}\text{C}$ (POCC) are noticeable, leading unequivocally to ^{13}C enriched signature when the water stored in floodplains rejoins the main channel (**Figure 6c**). The presumed influence of aquatic grasses, whose isotopic signature is heavy (-13 ‰, according to Victoria *et al.*, 1992), on the isotopic composition of POCC, seems to be confirmed here.

The amounts of DOM, both those observed at Óbidos and those reconstituted by modelling (model M6) exhibit a drastic decline of $[\text{DON}]$ and a correlative increase of $[\text{DOC}] / [\text{DON}]$ when the inundation of the floodplains occurs, i.e. during rising water stage. At this stage, the chemical nature of DOM, mobilized by surface runoff (Tardy *et al.*, 2005), is mainly soil-derived and refractory (Hedges *et al.*, 1986). Low C/N molecules, conveyors of positive charge (e.g. amino-acids) or hydrophobic (humic acids), are recognized to be the best candidates to sorption onto fine sediments (Aufdenkampe *et al.*, 2001), leading subsequently to rising C/N in the remaining DOM fraction. Conversely, the emptying of floodplains coincides with a rise of $[\text{DOC}]$ and $[\text{DON}]$ and a decrease of $[\text{DOC}]/[\text{DON}]$. This supports the idea that additional DOC and DON observed during the emptying of floodplains is autochthonous, is released by the decay of aquatic biomass which contains high proportion of amino acids (low C/N) compared to soil-derived DOM (Hedges *et al.*, 1994). The high variability of C/N tends to confirm that along a complete hydrological cycle, distinct pools of molecules (allochthonous soil-derived vs. autochthonous

river-derived), exhibiting contrasted reactivity and very dissimilar elemental composition, are exported by the Amazonian rivers (Amon & Benner, 1996a; Amon & Benner, 1996b).

5. Summary and concluding remarks

The six hydrochemical models which were tested provide a valuable insight on the main factors (hydrological source, water budget of the floodplains, nature of hydrobiological pattern i.e. photosynthesis vs. mineralization, air-water gaseous exchanges, etc) controlling the biogeochemical and sedimentary budgets of the Amazonian floodplains. The influence of floodplain and additional flow (small rivers, alluvial groundwater, direct precipitation) could be set in evidence for most of studied parameters (Bustillo, 2007). Unfortunately, due to lack of reliable data concerning the water chemistry of small tributaries, the influence of variable additional input (involving variable contribution of small rivers and alluvial groundwaters to the river flow) cannot be distinguished from the effects of the diagenesis operating in the floodplains. At the light of the results provided by the six mixing models, 3 main issues dealing with the biogeochemistry and hydrology of the floodplains are addressed:

- (1)- Coupling between sediment deposition and biogeochemical diagenesis;
- (2)- Organic metabolism of the river and its effects on the nature and intensity of biotic processes;
- (3)- Nature and intensity of abiotic processes, involving notably the sorption of DOM, the evaporation of wetlands and the river outgassing;

A companion paper (Bustillo *et al.*, submitted to *Earth Interactions*) aims to investigate more in detail these 3 topics which are intrinsically related and which determine most of the biogeochemical budget relative to Amazonian floodplains.

The magnitude and polarity of water exchanges between the Amazon River and its floodplains strongly influences the sedimentary and chemical signals measured in the river waters. The floodplains constitute widespread sites where major biotic and abiotic processes affecting the dynamics of transiting materials occur: sedimentation, remobilization of sediments, organic matter decay, CO₂ outgassing, etc.

Unexpectedly, the chemical trends observed upstream are sometimes accentuated downstream, as shown by the model M1. It supports the idea that the processes operating downstream are of same nature than those occurring upstream, prolonging therefore the imprint given by upstream rivers to the chemical baseline. Due to the increase in floodplain size as we move downstream, the impact of floodplain filling and draining on the biogeochemical qualities of the water are therefore amplified downstream. Taking into account the additional contribution of

ungauged areas, using ad-hoc constant characteristics to close the river budget (model M2) does not improve the performance of modelling compared to the simplest possible model M1. It means that the composition of the “additional” flow is probably very variable, more particularly because the alluvial groundwaters draining unweathered sediments deposited alongside the Amazon River (high TDS) and the small tributaries draining thick-sandy soils (intensively leached, low TDS) do not contribute synchronously to the river flow (Bustillo, 2007) and exhibit very distinct chemical characteristics.

The mixing model M3 sets in evidence a decrease of compositional differences between hydrological reservoirs as we move downstream. This homogenization might be the result of the mixture of waters having resided more or less durably in the hydrographic network. The main differences between incoming and outgoing compositions is attributable to the baseflow RB , and to a lesser extent, to the delayed direct runoff Rl . This would be the result of in-stream biogeochemical processes: the aquatic photosynthesis impacts RB whose contribution to river flow is maximum during lowest waters stage *i.e.* when autotrophic regime prevails, while organic matter decay impacts more particularly Rl ($\uparrow CO_2$, $\downarrow pH$, $\downarrow O_2$, $\downarrow \delta^{13}C-DIC$, $\downarrow DON$, $\downarrow DOC$ et $\downarrow DOC/DON$) whose contribution is maximum when the emptying of floodplains (where heterotrophic regime prevails) occurs.

The model M4, also based on variable hydrological source, involves the hydrobiological index $I_{BIO}^j = [O_2]_k^j - [CO_2]_k^j$, used as a tracer of autotrophic vs. heterotrophic regime. This improves considerably the performances of the simulations, compared to M3. The model M4 enables to (1) identify the parameters significantly influenced by in-stream processes and (2) to determine their response depending on the nature and magnitude of hydrobiological regime. Globally speaking, the hydrobiological regime promotes large variations of pH and $[O_2]$ which have direct repercussions on the biodynamics of other chemical variables. The autotrophic regime is dominant ($I_{BIO} > 0$) during lowest waters stage, when (1) the river turbidity is minimum, (2) when the river-floodplain connectivity is interrupted and (3) when the rate of incoming solar radiation reaching the water body is maximum (flow concentrated within the well-exposed main channel). The rises of pH and $[O_2]$ directly induced by aquatic photosynthesis coincide with losses of organic nitrogen to the benefit of mineral nitrogen, increase of $^{13}C/^{12}C$ for DIC (isotopic fractionation induced by aquatic photosynthesis), and losses of Ca^{2+} and HCO_3^- . The heterotrophic regime is dominant over the annual cycle, except during lowest waters stage. The heterotrophic signal is hugely amplified when the waters stored in the floodplains rejoin the main channel. Falling waters constitute privileged moments to appreciate the biogeodynamics of the floodplains, because their discharge in the main channel is intermittent. The decrease of pH and $[O_2]$ related to the heterotrophic regime coincide with

increase of $[\text{CO}_2]$, decrease of $\delta^{13}\text{C}(\text{DIC})$, decrease of $[\text{DOC}]/[\text{DON}]$ and rise of $\delta^{13}\text{C}(\text{POCC})$ which is interpreted as the result of the sequential release of autochthonous carbon (C_4 aquatic grasses).

The models M5 and M6 enable to test more specifically the impact of floodplains water balance (filling vs. emptying) on the differences of chemical concentrations between the tributaries and the Amazon River. This test appears to be very conclusive, setting in evidence that chemical signals observed in the Amazon River waters are thoroughly influenced by the magnitude and polarity of water exchanges between the Amazon River main channel and its floodplains.

List of parameters

Indexes: i stands for chemical species, j for the sampling station and k for the number of the sample.

Hydroclimatic features

RS : forwarded direct runoff; Ri : delayed direct runoff; RB : baseflow

QK : discharge of each individual runoff with K standing for RS , Ri or RB .

Geochemical characteristics (see Tardy *et al.*, 2004, for full details)

$Re = \text{SiO}_2 / \text{Al}_2\text{O}_3$ in altered products. Stoichiometry of clays formed by chemical weathering.

$\text{CO}_2 \text{ CARB}$ = CO_2 consumed by the alteration of carbonated rocks.

$\text{CO}_2 \text{ SIL}$ = CO_2 consumed by the alteration of crystalline rocks.

$\text{CO}_2 \text{ TOT} = \text{CO}_2 \text{ SIL} + 2.\text{CO}_2 \text{ CARB}$. DIC released by geochemical alteration.

$F_R = \text{CO}_2 \text{ CARB} / \text{CO}_2 \text{ TOT}$. Lithological index = part of DIC originating from the dissolution of carbonates.

WR = Chemical weathering rate ($\text{m}.\text{Ma}^{-1}$).

$F \text{ CO}_2$ = Rate of CO_2 consumption ($\text{TC}.\text{km}^{-2}.\text{a}^{-1}$).

Hydrochemical modelling

$I_{BIO}^j_k$ = Hydrobiological index based on the gaseous composition of the river water.

$K_{BIO}^j_i$: rate of uptake or release of chemical species associated to hydrobiological path.

$(\Delta Q_t / Q_t)_k^j$ = index of hydrograph stage (>0 during rising water, <0 during falling water and $=0$ for highest waters and lowest waters).

$\Delta Q_t^j = Q_t^j / Q_{tot,k}^j - 1$ = water balance of floodplains (<0 if filling and >0 if emptying).

α_i^j : variation of concentration in the river water associated to the variation of QR_s/Q_t in the main channel.

β_i^j : variation of concentration in the river water associated to the variation of QR_l/Q_t in the main channel.

γ_i^j : variation of concentration in the river water associated to the water balance of floodplains ΔQ_t^j .

δ_i^j : residual variation of concentration in the river water, for $\Delta QR_s/Q_t = \Delta QR_l/Q_t = \Delta Q_t = 0$.

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TABLE 1. Modelling of the Amazon river composition along a 2000-km longitudinal profile. Main principles and rules of calculations of 6 distinct mixing models: M1 and M2 are based on the variable contribution of regional sources to water and biogeochemical budgets; M3 and M4 rely on the variable contribution of hydrological sources (namely forwarded direct runoff: RS , delayed direct runoff: RI and baseflow: RB) with a correction for M4 taking into account the influence of the river processes; M5 and M6 are composite models taking into account the combined effects of the variable contributions relative to the regional and hydrological sources.

Model	Principle	Rule of calculation
M1	Variable contribution of major tributaries.	$C_{ijk} = \sum \sum \sum [C_{ijk,t} \times Q_{(jk,t)} / Q_{t(jk)}]$
M2	Step 1 : Variable contribution of major tributaries. Step 2 : mean concentration of the additional flow (small tributaries and alluvial aquifers).	$C_{ijk} = C_{ijk,M1} \times Q_{in(jk)} / Q_{o(jk)} + C_{ij(v)} \times [Q_{o(jk)} - Q_{in(jk)}] / Q_{o(jk)}$
M3	End-member mixing models : 3 reservoirs RS : forwarded direct runoff; RI : delayed direct runoff; RB : baseflow	$C_{ijk} = C_{ij(RS)} \times Q_{RS(jk)} / Q_{t(jk)} + C_{ij(RI)} \times Q_{RI(jk)} / Q_{t(jk)} + C_{ij(RB)} \times Q_{RB(jk)} / Q_{t(jk)}$
M4	Step 1 : Variable contribution of major tributaries. Step 2 : End-member mixing models taking into account biotic processes.	$C_{ijk(RB)} = C_{ij(RB),0} + K_{BIO(ij)} \times I_{BIO(jk)}$
M5	Step 1 : Variable contribution of major tributaries Step 2 : End-member mixing models calibrated on the relative differences (ΔC_{ijk}).	$\Delta C_{ijk} = f(3 \text{ covariates})$ $\Delta(QRS / Qt), \Delta(QRI / Qt) \text{ \& } \Delta Qt(I-O)$
M6	Step 1 : Variable contribution of major tributaries Step 2 : End-member mixing models calibrated on the relative differences (ΔC_{ijk}).	$\Delta C_{ijk} = f(2 \text{ covariates})$ (1)- River discharge Qt (2)-Default of water balance $\Delta Qt(I-O)$

i is the index of the chemical species ($ni = 44$), j is the index of the monitoring station ($nj = 10$) and k is the index of the sample ($nk = 8$).

TABLE 2. Mean parameters of the linear equation (α , β , r^2 , Bias) relating incoming and outgoing concentration (model M1) for 44 chemical parameters at 10 sampling stations (j) of the Amazon main stem. Data are presented for each chemical parameter (index i) and correspond to the mean values from 10 equations: $C_i^j(obs.) = \alpha_i^j \times C_i^j(calc.) + \beta_i^j$ ($n_j = 10$). The mean biases and r^2 are also given.

		Cations (μmolL^{-1})				μeqL^{-1}		Anions (μmolL^{-1})						mgL^{-1}	μmolL^{-1}		
k	pH	Na ⁺	K ⁺	Ca ²⁺	Mg ²⁺	S ⁺	S ⁻	HCO ₃ ⁻	Cl ⁻	NO ₃ ⁻	DOC ⁻	SO ₄ ²⁻	HPO ₄ ²⁻	DOC	DIC	CO ₂	O ₂
α	1.07	1.04	1.35	0.80	1.12	0.84	0.84	0.80	0.70	1.30	0.86	0.40	0.79	0.86	0.81	1.17	0.82
β	-0.59	-3	-9	25	-9	70	71	78	24	-2.6	1.9	23.4	0.23	0.35	98	-9	20
r^2	0.77	0.63	0.81	0.91	0.84	0.87	0.87	0.78	0.66	0.84	0.84	0.27	0.36	0.84	0.80	0.82	0.74
Bias	-0.08	-3	0	-37	-3	-84	-84	-67	-11	0.3	-1.3	-2.7	0.05	-0.24	-54	12	-12
Ave	6.99	144	27	273	55	827	827	618	92	11.2	20.1	42.1	0.76	3.76	759	141	159

	μmolL^{-1}	mgL^{-1}						$\delta^{13}\text{C}$			$\delta^{18}\text{O}$	C/N			Organic N (μmolL^{-1})		
k	SiO ₂	FSS	CSS	TSS	POCF	POCC	POC	DIC	POCF	POCC	H ₂ O	POCF	POCC	DOC	PONF	PONC	DON
α	0.89	0.86	0.26	0.76	0.88	0.24	0.77	1.14	0.99	0.61	0.95	0.41	0.60	0.69	0.72	0.26	0.86
β	15	8	36	36	0.05	0.38	0.1	1.4	-0.6	-10.9	-0.12	6.6	11.1	24.6	3.7	1.2	1.9
r^2	0.82	0.91	0.39	0.83	0.79	0.10	0.68	0.92	0.69	0.50	0.98	0.34	0.21	0.21	0.81	0.15	0.49
Bias	1	-26	-13	-38	-0.25	-0.08	-0.22	-0.57	-0.28	-0.19	0.23	0.35	1.63	-1.50	-2.38	-0.37	0.00
Ave	136	227	59	286	2.59	0.53	3.11	-14.32	-27.19	-27.72	-6.10	10.97	24.91	28.26	19.80	1.83	11.17

	Silicates (μmolL^{-1})				Carbonates (μmolL^{-1})		Carbon cycle (μmolL^{-1})			Add. Indexes	
k	NaSil	KSil	CaSil	MgSil	CaCO ₃	Dolomite	CO ₂ carb	CO ₂ sil	CO ₂ tot	FR	Re
α	1.67	1.35	1.67	1.67	0.76	1.45	0.96	1.29	0.80	1.93	0.34
β	-21	-8.54	-8	-7	13.55	-25.05	-31.63	-16.44	85.06	-0.40	1.53
r^2	0.34	0.81	0.34	0.34	0.67	0.80	0.79	0.26	0.79	0.64	0.14
Bias	8.0	0.0	3.2	2.6	-31.32	-5.73	-42.78	19.56	-65.99	-0.03	0.19
Ave	52.0	26.60	20.8	16.6	172.17	38.26	248.69	153.51	650.90	0.38	2.23

TABLE 3. Compared composition of the three individual runoff *Rs*, *Ri* and *Rb* for the 8 major tributaries (8 Rios, calculated data from discharge-weighing of the contributing reservoirs), and the stations of Paurá and Óbidos (data established from measured concentrations by multilinear regressions).

	k	mm/a		Cations (μmolL^{-1})				μeqL^{-1}		Anions (μmolL^{-1})					
		QK	pH	Na ⁺	K ⁺	Ca ²⁺	Mg ²⁺	S ⁺	S ⁻	HCO ₃ ⁻	Cl ⁻	NO ₃ ⁻	DOC ⁻	SO ₄ ²⁻	HPO ₄ ²⁻
Paurá	<i>Rs</i>	358	6.66	28	39	250	69	706	706	584	52	10.0	26	17	0.02
	<i>Ri</i>	490	6.34	33	21	102	21	302	302	232	17	0.6	28	12	0.41
	<i>Rb</i>	317	7.66	308	17	205	57	848	848	535	134	24.6	8	71	1.56
	Ave	1165	6.80	106	25	176	46	575	575	423	60	10.0	22	30	0.60
	R ²		0.60	1.00	0.31	0.55	0.71	0.68	0.68	0.59	0.85	0.80	0.49	0.25	0.90
Óbidos	<i>Rs</i>	345	6.63	28	44	246	68	703	703	586	68	11.4	22	8	0.53
	<i>Ri</i>	437	6.23	33	18	102	25	306	306	254	19	-2.2	25	5	0.68
	<i>Rb</i>	341	7.52	285	19	188	51	782	782	476	99	27.4	14	80	0.98
	Ave	1122	6.75	108	26	172	46	573	573	423	58	11.0	21	28	0.73
	R ²		0.48	1.00	0.16	0.26	0.28	0.36	0.36	0.33	0.69	0.72	0.17	0.36	0.05
8 Rios	<i>Rs</i>	344	6.59	28	41	302	57	789	790	628	3	3.9	41	57	0.07
	<i>Ri</i>	433	6.36	33	22	120	25	351	350	284	10	0.4	39	8	0.30
	<i>Rb</i>	340	7.81	280	11	185	62	788	787	490	196	25.2	-10	41	1.36
	Ave	1117	6.87	107	25	196	46	619	618	453	65	9.0	25	33	0.55
	R ²		0.73	1.00	0.69	0.80	0.90	0.85	0.85	0.75	0.98	0.82	0.76	0.83	0.46

	k	mm/a	mgL ⁻¹	μmolL^{-1}				mgL ⁻¹						$\delta^{18}\text{O}$
		QK	DOC	DIC	CO ₂	O ₂	SiO ₂	FSS	CSS	TSS	POC _F	POC _C	POC	H ₂ O
Paurá	<i>Rs</i>	358	4.9	820	237	116	123	620	58	679	5.77	0.57	6.34	-7.8
	<i>Ri</i>	490	5.3	466	236	116	107	20	51	70	1.17	0.52	1.69	-7.3
	<i>Rb</i>	317	1.5	494	-44	250	154	0	-24	-23	-0.43	-0.18	-0.61	-0.9
	Ave	1165	4.1	583	160	153	125	199	33	232	2.15	0.35	2.49	-5.7
	R ²		0.49	0.59	0.50	0.48	0.44	1.00	0.40	0.99	0.94	0.55	0.93	0.46
Óbidos	<i>Rs</i>	345	4.2	802	216	120	117	610	63	673	5.3	0.8	6.0	-7.1
	<i>Ri</i>	437	4.9	519	264	90	128	15	59	73	2.0	0.4	2.4	-7.6
	<i>Rb</i>	341	2.7	445	-30	262	144	0	-32	-31	-0.3	-0.3	-0.5	-1.6
	Ave	1122	4.0	583	160	151	129	193	33	226	2.3	0.3	2.6	-5.6
	R ²		0.17	0.43	0.49	0.67	0.17	1.00	0.52	0.99	0.85	0.85	0.89	0.39
8 Rios	<i>Rs</i>	344	7.7	877	248	76	95	665	192	857	5.5	1.4	7.0	-10.3
	<i>Ri</i>	433	7.4	522	234	105	101	20	48	68	1.0	0.5	1.5	-8.4
	<i>Rb</i>	340	-2.0	389	-95	320	191	0	-94	-93	0.5	-0.6	-0.1	1.7
	Ave	1117	4.6	591	138	162	127	213	49	262	2.2	0.4	2.7	-5.9
	R ²		0.78	0.67	0.59	0.39	0.55	1.00	0.81	0.99	0.95	0.72	0.94	0.41

	k	mm/a	$\delta^{13}\text{C}$			C/N			Carbon cycle (μmolL^{-1})						
		QK	DIC	POC _F	POC _C	POC _F	POC _C	DOC	DIC	POC _F	POC _C	PON _F	PON _C	DOC	DON
Paurá	<i>Rs</i>	358	-14.5	-25.7	-28.5	7.5	28.8	22.9	810	480	48	55.3	2.15	404	8.5
	<i>Ri</i>	490	-18.7	-27.8	-26.9	11.2	22.0	-10.1	461	98	44	6.5	2.00	438	25.1
	<i>Rb</i>	317	-11.6	-28.0	-28.0	9.9	11.6	109.0	490	-36	-15	-4.7	-0.51	125	-9.5
	Ave	1165	-15.50	-27.22	-27.70	9.7	21.2	32.5	576	179	29	18.4	1.36	343	10.6
	R ²		0.67	0.61	0.68	0.51	0.11	0.81	0.60	0.94	0.55	0.91	0.30	0.49	0.67
Óbidos	<i>Rs</i>	345	-14.6	-27.4	-28.3	10.5	21.9	48.9	812	439	65	45.0	3.1	349	5.3
	<i>Ri</i>	437	-19.3	-28.1	-27.5	12.9	17.2	3.6	525	169	35	11.7	1.9	408	24.3
	<i>Rb</i>	341	-11.8	-26.1	-28.4	6.1	21.0	35.3	450	-22	-23	2.7	-1.1	224	4.1
	Ave	1122	-15.56	-27.26	-28.01	10.1	19.8	27.1	590	194	26	19.2	1.3	334	12.3
	R ²		0.46	0.29	0.17	0.57	0.08	0.29	0.43	0.85	0.85	0.76	0.79	0.17	0.41
8 Rios	<i>Rs</i>	344	-17.0	-26.4	-25.7	6.8	24.4	41.1	882	461	119	54.2	5.2	642	15.6
	<i>Ri</i>	433	-18.4	-27.4	-26.9	10.2	26.5	24.5	525	79	42	7.5	1.6	611	22.9
	<i>Rb</i>	340	-8.2	-27.1	-29.9	11.3	16.1	22.3	392	41	-52	-0.1	-1.9	-163	-1.3
	Ave	1117	-14.86	-27.02	-27.47	9.5	22.7	29.0	594	185	37	19.6	1.6	385	13.3
	R ²		0.78	0.50	0.10	0.34	0.32	0.42	0.67	0.95	0.72	0.97	0.74	0.76	0.47

	k	mm/a	Sil (μmolL^{-1})			Carb (μmolL^{-1})		CO ₂ cycle (μmolL^{-1})				
		QK	NaSil	CaSil	MgSil	CaCO ₃	Dolom.	CO ₂ CARB	CO ₂ SIL	CO ₂ TOT	FR	Re
Paurá	<i>Rs</i>	358	-25	-10	-8	166	77	319	-21	619	0.56	1.45
	<i>Ri</i>	490	16	7	5	68	16	100	61	261	0.37	1.73
	<i>Rb</i>	317	174	70	56	62	2	66	440	571	0.07	3.89
	Ave	1165	46.7	18.7	14.9	96	31	158	139	456	0.35	2.23
	R ²		0.90	0.90	0.90	0.41	0.63	0.56	0.94	0.61	0.75	0.90
Óbidos	<i>Rs</i>	345	-40	-16	-13	175	80	335	-53	621	0.60	1.46
	<i>Ri</i>	437	14	6	4	71	21	113	52	278	0.38	1.44
	<i>Rb</i>	341	187	75	60	40	-7	26	473	519	0.02	4.12
	Ave	1122	49.8	19.9	15.9	94	30	154	148	456	0.34	2.26
	R ²		0.91	0.91	0.91	0.50	0.39	0.45	0.97	0.35	0.78	0.76
8 Rios	<i>Rs</i>	344	25	15	10	185	49	284	101	669	0.45	2.54
	<i>Ri</i>	433	23	14	9	85	17	120	78	319	0.37	1.86
	<i>Rb</i>	340	84	49	33	75	35	146	215	508	0.28	1.81
	Ave	1117	42.0	24.7	16.8	113	33	179	127	484	0.37	2.06
	R ²		0.89	0.89	0.89	0.57	0.86	0.69	0.88	0.74	0.44	0.53

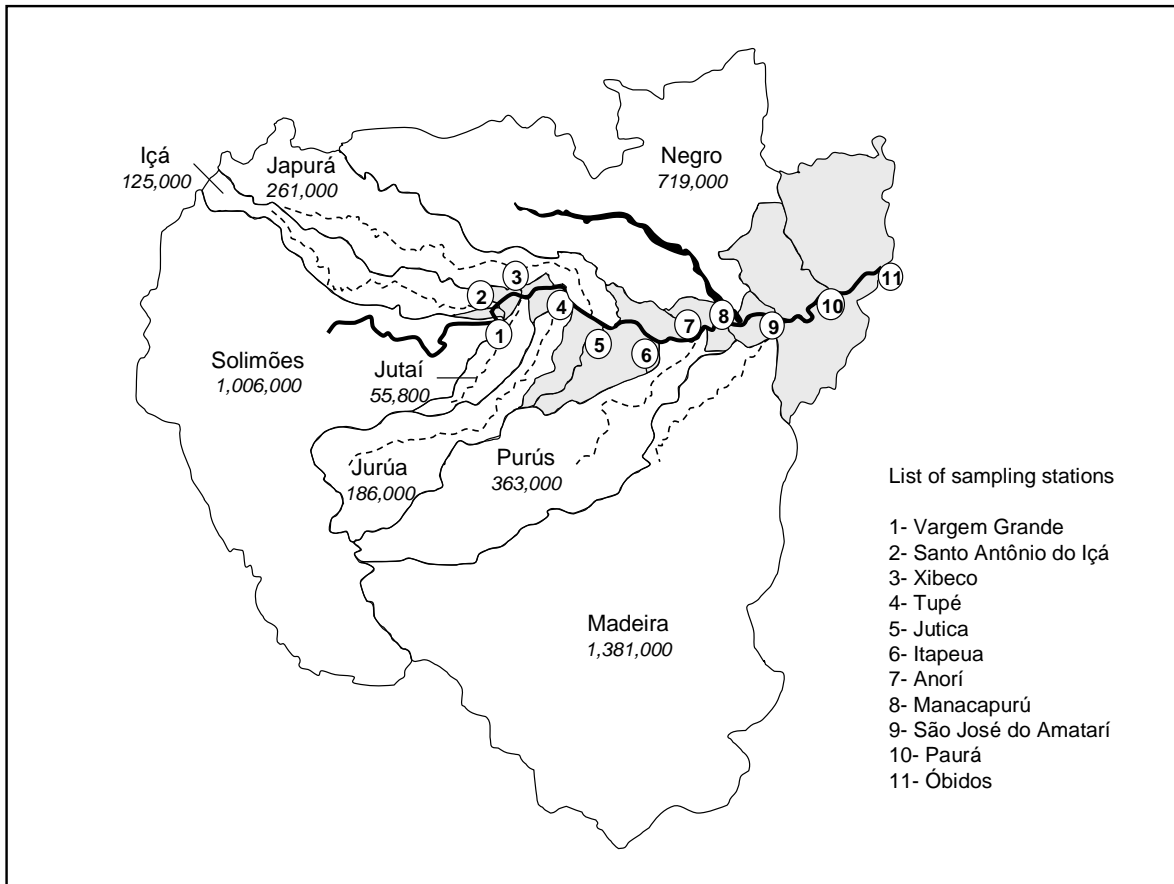


FIG. 1. Map of the Amazon basin upstream from Óbidos showing the major tributaries and the geographical repartition of small tributaries (areas coloured in grey) along the Amazon River main stem. Numbers in italics stand for the drainage area of major sub-basins, expressed in km².

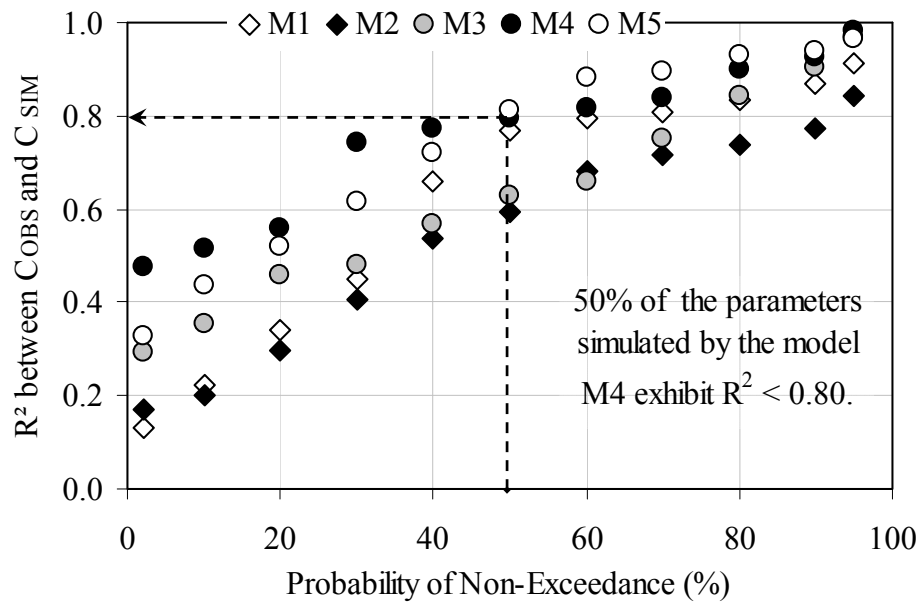


FIG. 2. Compared performance of the five hydrochemical models (M1 to M5) based on the Probability of Non-Exceedance of the determination coefficient R^2 established by confronting simulated and observed concentrations (or isotopic values). Data represented are obtained by gathering the results of 42 chemical parameters for the station of Óbidos, outlet of the studied reach.

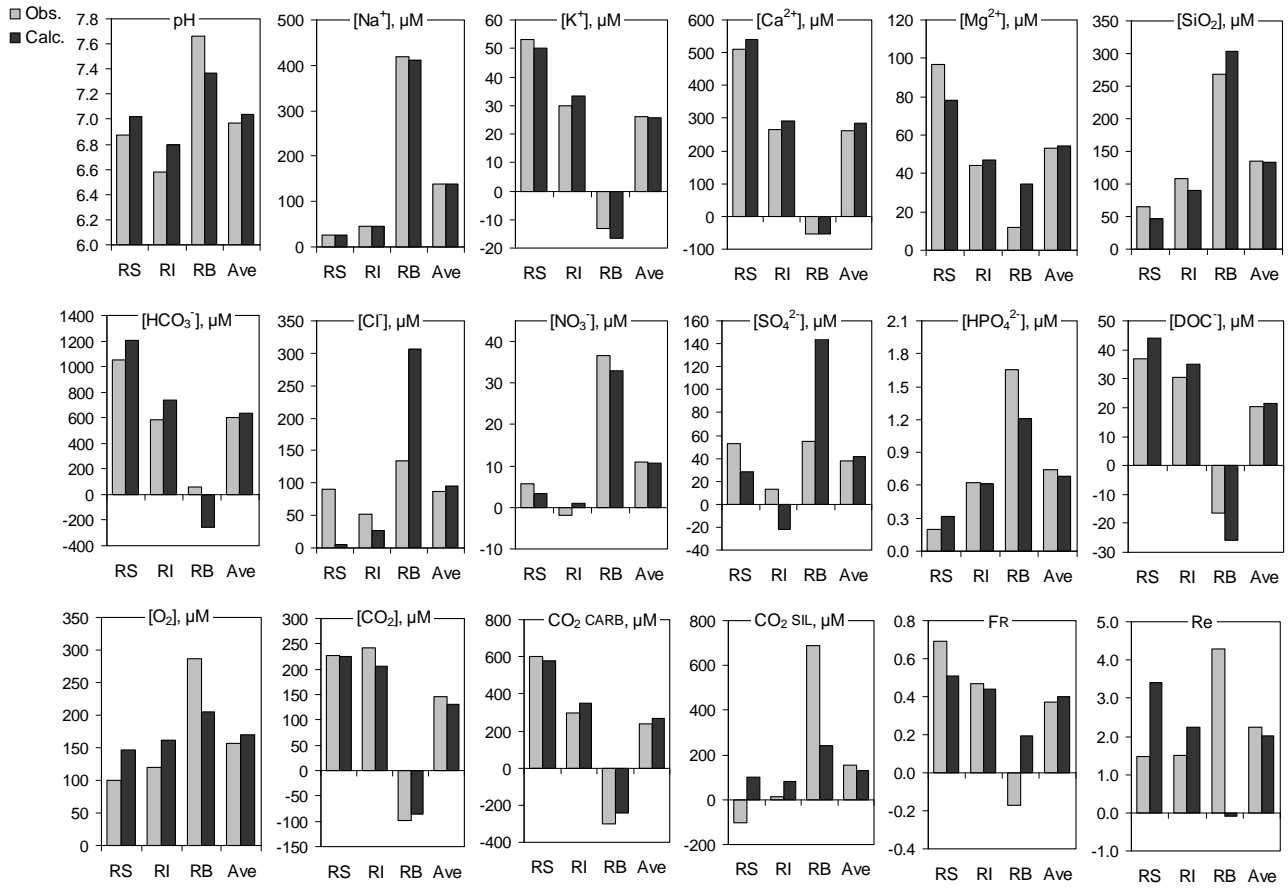


FIG. 3a. Composition of (i) the three individual runoffs *RS*, *RI* and *RB* and (ii) of the river (*Ave*) obtained by averaging the model M3's outcomes of 9 stations located on the studied Amazonian reach. Calculated data (noted Calc., see Eq.16) resulting from the discharge-weighting of runoffs composition of the major tributaries, are compared to observed data (noted Obs.), obtained by multilinear regression. Dissolved species and biogeochemical indices (see list of parameters).

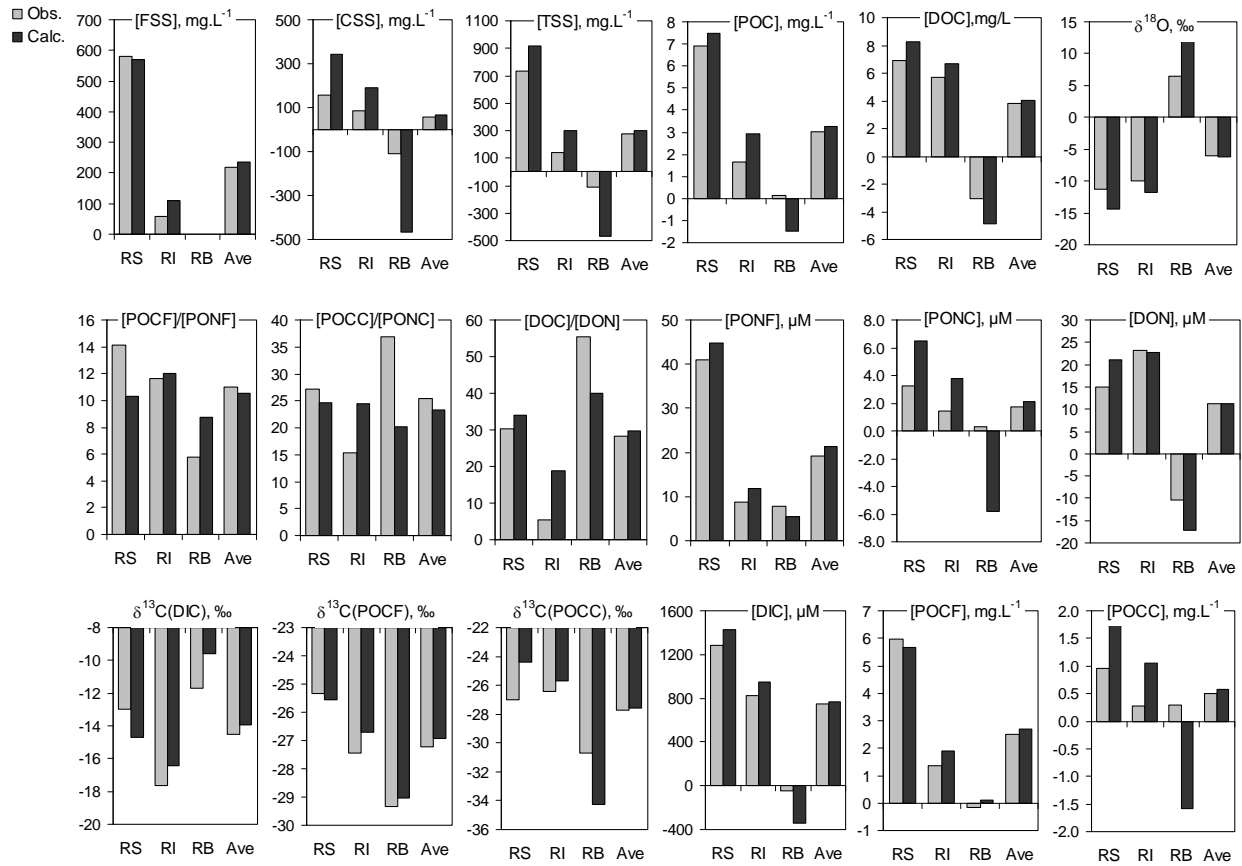


FIG. 3b. Composition of (i) the three individual runoffs *RS*, *RI* and *RB* and (ii) of the river (*Ave*) obtained by averaging the model M3's outcomes of 9 stations located on the studied Amazonian reach. Calculated data (noted Calc., see Eq.16), resulting from the discharge-weighting of runoffs composition of the major tributaries, are compared to observed data (noted Obs.), obtained by multilinear regression. Suspended sediments, organic carbon and nitrogen, C/N molar ratios, and isotopic signature of carbon ($\delta^{13}\text{C}$) and water ($\delta^{18}\text{O}$).

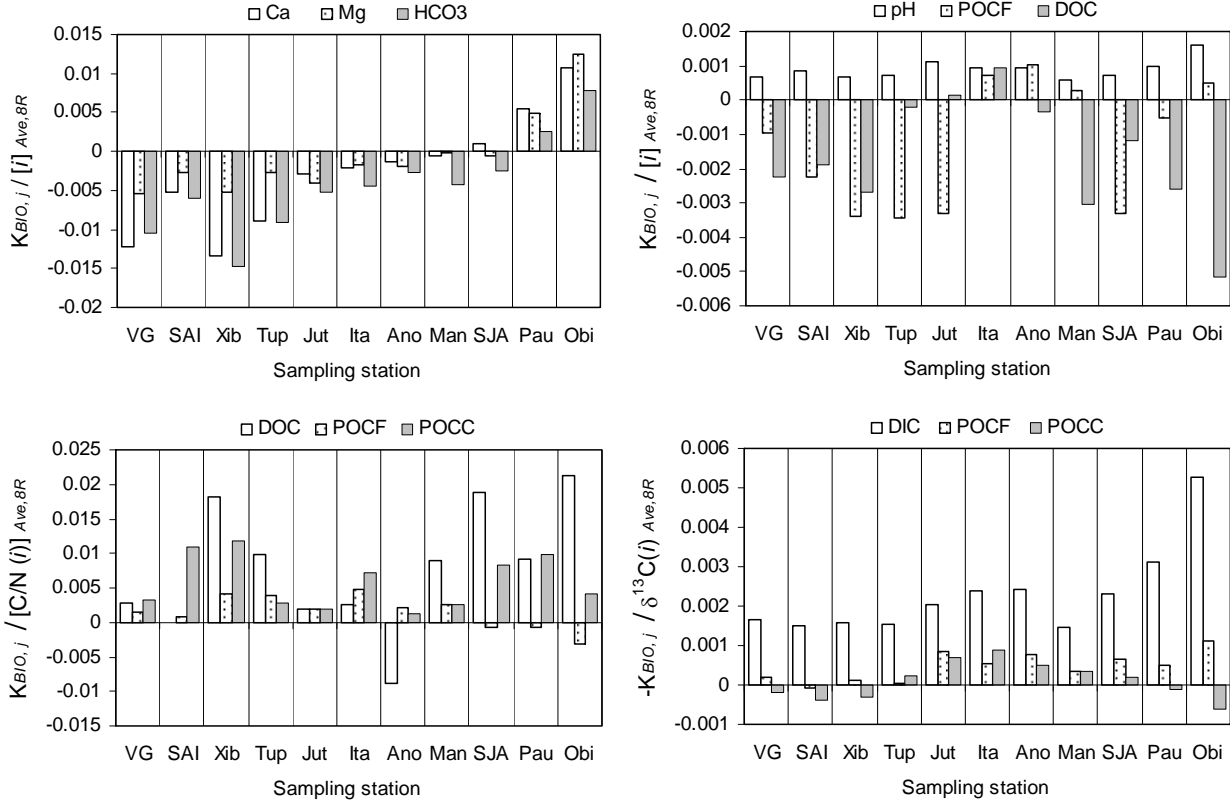


FIG. 4. Influence of the hydrobiological regime, appreciated by $K_{BIO,i}^j / [i]_{Ave,8R}$ for 12 chemical parameters (noted i) at the 11 sampling stations (noted j) over the Amazon River longitudinal profile: Vargem Grande (VG), Santo Antonio do Iça (SAI), Xibeco (Xib), Tupe (Tup), Jutica (Jut), Itapeua (Ita), Anori (Ano), Manacapuru (Man), São Jose da Amajari (SJA), Paura (Pau) and Óbidos (Óbi) the outlet of the studied reach.

$[i]_{Ave,8R}$ is the mean concentration of i , calculated by discharge weighing the inputs of the 8 major tributaries upstream from Obidos; $K_{BIO,i}^j$ is a calibrated parameter (outcomes of the model M4, see Eq.11) corresponding to the rate of uptake or release of each bioactive element (i) for each station (j) associated to biologically-mediated processes in the river water and describing thus the response of chemical parameters to $I_{BIO,k}^j = [O_2]_k^j - [CO_2]_k^j$.

$K_{BIO,i}^j > 0$ means that C_i^j rises with photosynthetical pathways ($I_{BIO,k}^j > 0$) and decreases with mineralization pathways ($I_{BIO,k}^j < 0$), and vice versa.

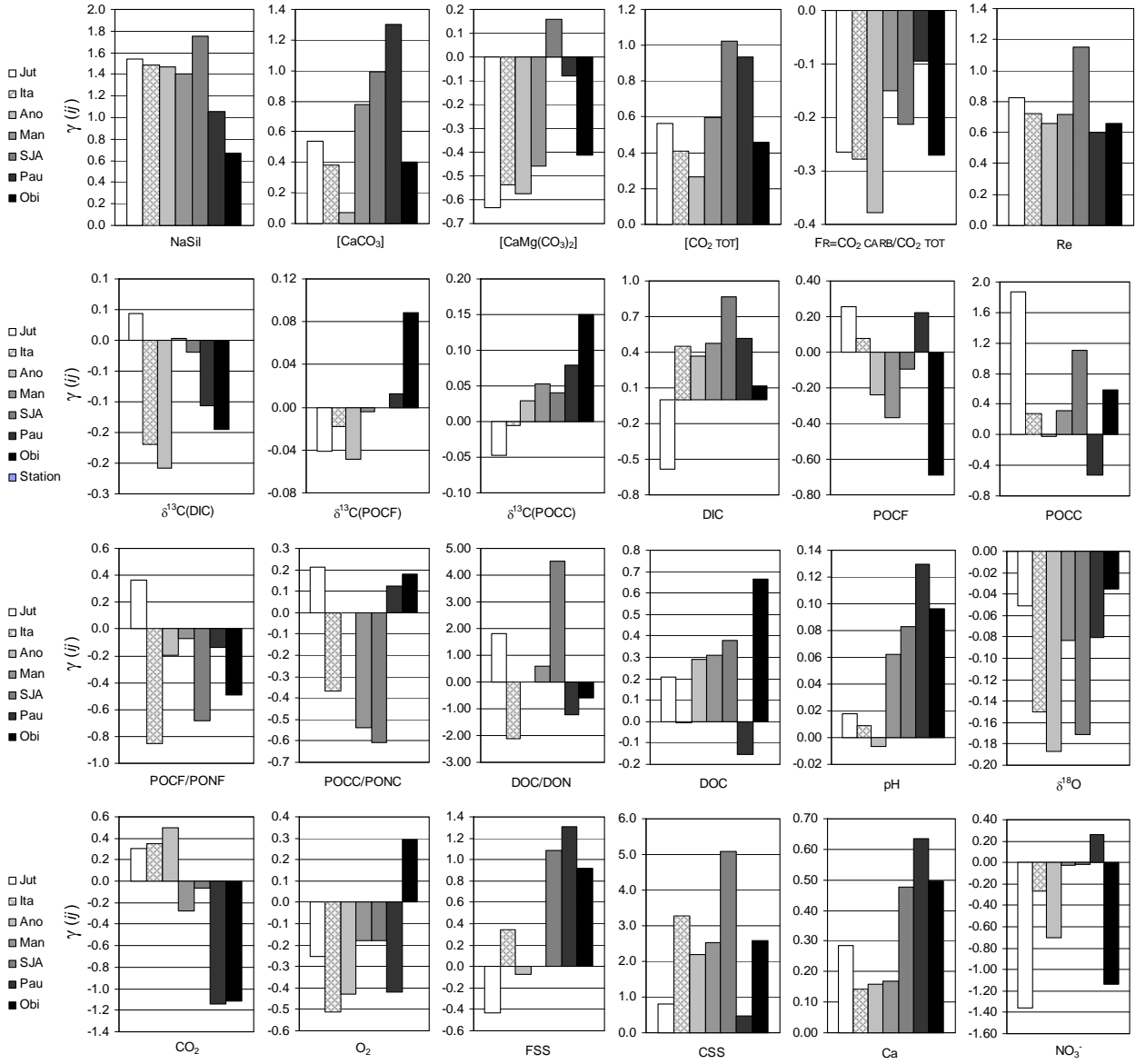


FIG. 5. Variation of the coefficients γ_i^j over the Amazon River (longitudinal profile between Jutica and Óbidos) obtained by the model M5 for a sample of 24 chemical parameters (i) and 7 sampling stations (j): Jutica (Jut), Itapeua (Ita), Anori (Ano), Manacapuru (Man), São Jose da Amatari (SJA), Paura (Pau) and Óbidos (Óbi) the outlet of the studied reach. The coefficients γ_i^j enable to track the influence of floodplains water balance on the compositional changes of water chemistry for a given parameter (i) at a given station (j): $\gamma_i^j > 0$ indicates that concentrations are higher (all other things being equal) when the floodplains drain [$\Delta Q t_k^j > 0$] and vice versa. For example, $\gamma_i^j = 0.92$ for fine suspended sediments (FSS) at Óbidos, indicating that [FSS] in the outgoing flow increases by 92% compared to [FSS] in the incoming flow (data calculated by discharge-weighting chemical signals from the 8 tributaries upstream from Óbidos) when $\Delta Q t_k^j = 100\%$ i.e. outflow = $2 \times$ inflow.

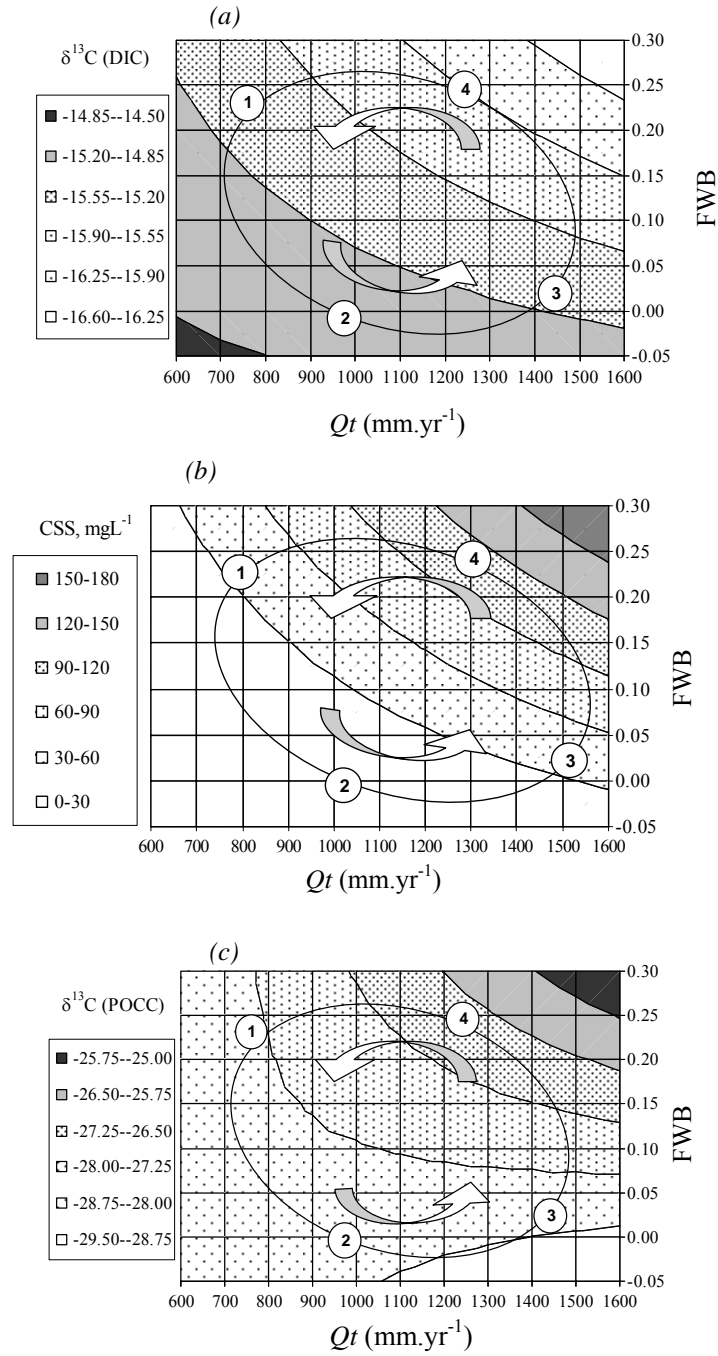


FIG. 6. Mean simulated variations of (a) $\delta^{13}C$ (DIC), (b) [CSS] and (c) $\delta^{13}C$ (POCC) as a function of the river outflow (Qt) and the floodplain water balance $FWB = \Delta Qt_k^j = Qt_k^{j,obs} / Qt_{tot,k}^j - 1$. The fluctuations modelled by M6 over an annual cycle, at the station of Óbidos, are represented by arrows, setting in evidence: (a) $^{13}C/^{12}C$ depletion during falling waters (path 3 → 4); (b) sedimentation patterns on the 1 → 2 → 3 paths and remobilization patterns on the 3 → 4 → 1 paths; (c) the exportation of the várzeas grasses ($^{13}C/^{12}C$ enriched) towards the main channel during the falling water stage (3 → 4 path). The hydrological sequence is: (1) lowest waters with outflow =

inflow (FBW = 0) → (2) rising waters, with outflow < inflow (FBW < 0) → (3) highest waters, with outflow = inflow (FBW = 0) → (4) falling waters, outflow > inflow (FBW > 0) → (1) lowest waters.